

## AUGMENTING A SEA OF DATA WITH DYNAMICS: OCEAN STATE & PARAMETER ESTIMATION

Patrick Heimbach Nora Loose, An T. Nguyen, Helen Pillar, Tim Smith, ..., CRIOS & ECCO groups *The University of Texas at Austin, TX* 

> https://crios-ut.github.io https://ecco-group.org

## Overview

- 1. Oceanography: a sparse-data science
- 2. ECCO: deterministic inverse problem in a statistical/Bayesian framework
- 3. DA filters versus smoothers
- 4. Science applications: time-evolving property budgets
- 5. Beyond optimization: adjoints for dynamical attribution
- 6. Hessian uncertainty quantification & observing system design

## **Oceanography:** A sparse data problem ...

**Observational sampling** coverage for ocean temperature in the upper 2000 m 1950 - 2010 (mean ocean depth: ~ 3900 m)





(colors refer to depth ranges)

An eclectic observing system with disparate sensors















# Two incomplete knowledge reservoirs

an eclectic, patchy
 global ocean observing system

#### • numerical models

that require uncertain inputs

## Two incomplete knowledge reservoirs

- Can we <u>optimally combine</u> these two knowledge reservoirs?
- Can we do so in a manner that provides <u>useful for climate analysis</u>?
  - causal, dynamical attribution
  - detecting small, residual signals
- Can we provide <u>measures of uncertainties</u> with these?
- Can we use simulation to <u>inform efficient</u> <u>observing strategies</u>?

## ECCO: Estimating the Circulation and Climate of the Ocean

A <u>multi-platform</u>, <u>multi-instrument</u> synthesis effort that integrates ocean and marine ice observations with equations of motion (model)

















- Physical Oceanography (PO)
- Cryosphere
- Modelling, Analysis, and Prediction (MAP)
- Advancing Collaborative **Connections for Earth** System Science (ACCESS)

## Learning from sparse, heterogeneous observations and models



## What is Data Assimilation?

## Kaminski et al., The Cryosphere (2015): "Ideally, ...

- ... all observational data streams are interpreted simultaneously,
- ... with the process information provided by the model,
- ... [which leads to] a consistent picture of the state of the system,
- ... that balances all the observational constraints,
- ... taking into account all the respective uncertainty ranges."

#### Penny et al., Front. Mar. Sci. (2019):

"DA allows information provided from observations to be propagated in time and space to unobserved areas using the dynamical and physical constraints imposed by numerical models."

## Kaipio & Somersalo (2005) Statistical and Computational Inverse Problems

- Formulate inverse problem as statistical quest for information
- Given observable & unobservable quantities
- **Solve** inverse problem:
  - unobservable (or unobserved) quantities are of primary interest
  - extract information and assess uncertainty based on <u>all available knowledge</u> of measurement process, of models, and of prior information of unknowns
- Solution of inverse problem is a posterior PDF of unknowns

In the following, we adopt Isaac et al., J. Comp. Phys. (2015)

#### Incomplete knowledge reservoirs

Model (ocean GCM)

$$\mathbf{x}(t)\,=\,\mathcal{L}(\,\mathbf{x}(t-1)\,,\,\mathcal{B}\,\mathbf{u}(t-1)\,)$$

#### Observations

$$\mathbf{y}(t) = \mathcal{E}\,\mathbf{x}(t) + \mathbf{n}(t)$$

- $\mathbf{x}(t)$  : model state vector at time t
- $\mathbf{y}(t)$  : observation vector at time t
- $\mathbf{u}(t)$ : vector of uncertain (control) variables at time t
- $\mathbf{n}(t)$  : residual noise of model-data misfit at time t
  - $\mathcal{L}$  : nonlinear model operator
  - $\mathcal{E}$  : state-to-observable map
  - $\mathcal{B}$ : parameter-to-state map

▲□ ▶ ▲ □ ▶ ▲ □ ▶

< 🗆 🕨

T

SQA

#### Statistical inference over space of uncertain parameters u

 $\pi_{post}(\mathbf{u} \mid \mathbf{y})$ : posterior distribution of uncertain parameters  $\mathbf{u}$ , given observations  $\mathbf{y}$ 

- combines prior PDF  $\pi_{prior}(\mathbf{u})$  with
- likelihood PDF  $\pi_{like}(\mathbf{y} \mid \mathbf{u})$

#### Bayes' Theorem

$$\pi_{\textit{post}}(\mathbf{u}\,|\,\mathbf{y})\,\propto\,\pi_{\textit{prior}}(\mathbf{u})\,\pi_{\textit{like}}(\mathbf{y}\,|\,\mathbf{u})$$

ヘロッ ヘヨッ ヘヨッ

Expensive forward model & high-dimensional space of uncertain parameters necessitates assumptions:

- $\pi_{prior}(\mathbf{u})$  is Gaussian
- prior error covariance may impose smoothness, e.g., via elliptic PDE operator (inverse Laplacian)
- In difference between predicted observables,  $\mathcal{E} \mathbf{x}(t)$ , and actual observations,  $\mathbf{y}(t)$  captured by noise vector

$$\mathbf{n}(t) = \mathcal{E} \, \mathbf{x}(t) - \mathbf{y}(t)$$

which obeys Gaussian statistics

ヘロマ ヘロマ ヘロマ

#### Generalize to parameter-to-observable map

$$\mathcal{F}(\mathbf{u},t) \,=\, \mathcal{E}\circ\mathcal{L}(\,\mathbf{x}(t-1)\,,\,\mathcal{B}\,\mathbf{u}(t-1)\,)$$

with generally nonlinear  $\mathcal{F}$ , such that

$$\mathcal{E} \mathbf{x}(t) - \mathbf{y}(t) = \mathcal{F}(\mathbf{u}, t) - \mathbf{y}(t)$$

leads to ...



< ロ > < 回 > < 回 > < 回 > < 回 >

500

#### From statistical to deterministic inversion

Mean  $\mathbf{u}_{MAP}$  of posterior distribution  $\pi_{post}(\mathbf{u})$  is parameter vector maximizing  $\pi_{post}$ , called **maximum a posteriori** (MAP) point

- found by minimizing  $-\log(\pi_{post})$
- solve optimization / deterministic inverse problem

$$\mathbf{u}_{MAP} = \operatorname*{arg\,min}_{\mathbf{u}} \left\{ -\frac{1}{2} || \, \mathcal{F}(\mathbf{u}) - \mathbf{y} \, ||_{\Gamma_{obs}^{-1}} - \frac{1}{2} || \, \mathbf{u} - \mathbf{u}_{prior} \, ||_{\Gamma_{prior}^{-1}} \right\}$$

This is the ECCO parameter & state estimation problem

< □ > < □ > < □ > < □ >

#### Approximation of the posterior PDF by posterior error covariance $\Gamma_{post}$

Posterior error covariance given by inverse of the Hessian at  $\mathbf{u}_{MAP}$ 

$$\Gamma_{post} = \left[ H_{misfit}(\mathbf{u}_{MAP}) + \Gamma_{prior}^{-1} 
ight]^{-1}$$

where

$$H_{misfit}(\mathbf{u}_{MAP}) \approx F^T \Gamma_{obs}^{-1} F$$

Gauss-Newton Hessian approximation

with:

- F: tangent linear operator of  $\mathcal{F}$
- $F^{T}$ : adjoint operator of  $\mathcal{F}$

< ロ > < 回 > < 回 > < 回 > < 回 >

Consider "perfect" nonlinear model  $\mathcal{L}$ , and observations **y** with noise **n** 

$$J(\mathbf{u}) = \frac{1}{2} \sum_{t=1}^{t_f} \left[ \mathcal{E} \mathbf{x}(t) - \mathbf{y}(t) \right]^T \Gamma_{obs}^{-1} \left[ \mathcal{E} \mathbf{x}(t) - \mathbf{y}(t) \right]$$
$$+ \frac{1}{2} \sum_{t=0}^{t_f - 1} \left[ \mathbf{u}(t) - \mathbf{u}_{prior}(t) \right]^T \Gamma_{prior}^{-1} \left[ \mathbf{u}(t) - \mathbf{u}_{prior}(t) \right]$$
$$= J_{misfit} + J_{prior}$$

Extend to Lagrange function  $\mathcal{J}$ , introducing Lagrange multipliers  $\mu(t)$ :

$$\mathcal{J}(\mathbf{u}, oldsymbol{\mu}) \,=\, J(\mathbf{u}) \,-\, \sum_{t=1}^{t_f} oldsymbol{\mu}^{\mathcal{T}} \left[\, \mathbf{x}(t) \,-\, \mathcal{L}(\mathbf{x}(t-1))\,
ight]$$

Patrick Heimbach ECCO as a statistical/Bayesian inverse problem

200

#### Lagrange multiplier method:

Stationary point of  $\mathcal J$  leads to set of normal equations:

$$\begin{split} \frac{\partial \mathcal{J}}{\partial \boldsymbol{\mu}(t)} &= \boldsymbol{x}(t) - \mathcal{L}(\boldsymbol{x}(t-1)) = 0 & 1 \leq t \leq t_f \\ \frac{\partial \mathcal{J}}{\partial \boldsymbol{x}(t)} &= \frac{\partial J}{\partial \boldsymbol{x}(t)} - \boldsymbol{\mu}(t) + \left[\frac{\partial \mathcal{L}(\boldsymbol{x}(t))}{\partial \boldsymbol{x}(t)}\right]^T \boldsymbol{\mu}(t+1) = 0 & 0 < t < t_f \\ \frac{\partial \mathcal{J}}{\partial \boldsymbol{x}(t_f)} &= \frac{\partial J}{\partial \boldsymbol{x}(t_f)} - \boldsymbol{\mu}(t_f) = 0 & t = t_f \\ \frac{\partial \mathcal{J}}{\partial \boldsymbol{x}(0)} &= \frac{\partial J}{\partial \boldsymbol{x}(0)} - \left[\frac{\partial \mathcal{L}(\boldsymbol{x}(0))}{\partial \boldsymbol{x}(0)}\right]^T \boldsymbol{\mu}(1) & t_0 = 0 \end{split}$$

<ロ> <同> <同> <巨> <巨>

T

DQA

"Variational" hints that we need a gradient:

- gradient of J with respect to unknown/uncertan or control variables u
- Here: Vary initial conditions,  $\mathbf{x}(0)$  such as to minimize J

BUT: J depends not just on  $\mathbf{x}(0)$ , but on all  $\mathbf{x}(t)$ .

- consider nonlinear model  $\mathbf{x}(t+1) = \mathcal{L}(\mathbf{x}(t))$
- linearized operator is state transition matrix or Jacobian L

$$\delta x(t+1) = \frac{\partial x(t+1)}{\partial x(t)} \delta x(t) = \mathbf{L} \delta x(t)$$

Need chain rule of differentiation:

$$J = J(x_0, x_1, x_2, ..., x_{t_f})$$
  
=  $J(x_0, \mathcal{L}(x_0), \mathcal{L}(\mathcal{L}(x_0)), ..., \mathcal{L}^{N_{t_f}}(x_0))$ 

Patrick Heimbach ECCO as a statistical/Bayesian inverse problem

200

$$\begin{split} \mu_{0} &= \frac{\partial J}{\partial x_{0}} = \sum_{1 \leq t \leq t_{f}} \frac{\partial x_{t}}{\partial x_{0}} \left( \frac{\partial J}{\partial x_{t}} \right) \\ &= \frac{\partial x_{1}}{\partial x_{0}} \left( \frac{\partial J}{\partial x_{1}} \right) + \frac{\partial x_{1}}{\partial x_{0}} \frac{\partial x_{2}}{\partial x_{1}} \left( \frac{\partial J}{\partial x_{2}} \right) \\ &+ \ldots + \frac{\partial x_{1}}{\partial x_{0}} \cdots \frac{\partial x_{t_{f}}}{\partial x_{t_{f}-1}} \left( \frac{\partial J}{\partial x_{t_{f}}} \right) \\ &= \mathbf{L}^{T} \frac{\partial J}{\partial x_{1}} + \mathbf{L}^{T} \mathbf{L}^{T} \frac{\partial J}{\partial x_{2}} + \ldots + \mathbf{L}^{T} \cdots \mathbf{L}^{T} \frac{\partial J}{\partial x_{t_{f}}} \end{split}$$

 $L^{T}$ : is the adjoint model (and L is the tangent linear model)  $\mu_{t} = \left(\frac{\partial J}{\partial x_{t}}\right)$ : Lagrange multiplier or dual state at time t

< □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > < □ > <

For intermediate step of the adjoint model integration one obtains:

$$\mu_{t} = \frac{\partial J}{\partial x_{t}} = \mathbf{L}^{T} \frac{\partial J}{\partial x_{t+1}} + \mathbf{E}^{T} \Gamma_{obs}^{-1} \left[ \mathcal{E} x_{t} - y_{t} \right]$$
$$= \mathbf{L}^{T} \left( \mathbf{L}^{T} \frac{\partial J}{\partial x_{t+2}} + \mathbf{E}^{T} \Gamma_{obs}^{-1} \left[ \mathcal{E} x_{t+1} - y_{t+1} \right] \right)$$
$$+ \mathbf{E}^{T} \Gamma_{obs}^{-1} \left[ \mathcal{E} x_{t} - y_{t} \right]$$

- The adjoint model L<sup>T</sup> propagates µ<sub>t</sub> (the sensitivity of J with respect to all earlier states x<sub>t</sub>) backward in time to x<sub>0</sub>;
- Each model-data misfit (i.e. innovation vector  $\mathcal{E}x_t y_t$ ) is a source of sensitivity;
- The gradient of J with respect to x<sub>0</sub> takes into account (and weighs) the size of all misfit terms, all (inverse) error covariances, and all (linearized) model dynamics.

DQA

Make explicit time-varying "forcing" in model  $\mathcal{L}$ , with known part  $\mathbf{B} q_k$ , uncertain part  $\mathbf{\Gamma} u_t$ 

$$x_{t+1} = \mathcal{L}(x_t, \mathbf{B} q_t, \mathbf{\Gamma} u_t)$$

Least-squares estimation problem now is:

$$J(x) = \sum_{0 \le t \le t_f} \left\{ \begin{bmatrix} \mathcal{E}x_t - y_t \end{bmatrix}^T \mathbf{\Gamma}_{obs}^{-1} \begin{bmatrix} \mathcal{E}x_t - y_t \end{bmatrix} + \begin{bmatrix} u_t - u_{prior} \end{bmatrix}^T \mathbf{\Gamma}_{prior}^{-1} \begin{bmatrix} u_t - u_{prior} \end{bmatrix} \right\} + \begin{bmatrix} x_0 - x^b \end{bmatrix}^T \mathbf{\Gamma}_{init}^{-1} \begin{bmatrix} x_0 - x^b \end{bmatrix}$$

#### ▲□▶▲□▶▲□▶▲□▶ □ のへで

Patrick Heimbach ECCO as a statistical/Bayesian inverse problem

Lagrange function  $\mathcal{L}$  now takes on the form:

$$\mathcal{J} = J + \sum_{0 \leq t \leq t_f} \mu(t)^T \left[ \mathbf{x}(t+1) - \mathcal{L}(\mathbf{x}(t), \mathbf{B} \mathbf{q}(t), \mathbf{\Gamma}_{prior} \mathbf{u}(t)) \right]$$

And we have additional normal equation for gradient w.r.t.  $\mathbf{u}(t)$ 

$$\frac{\partial \mathcal{J}}{\partial \mathbf{x}(t)} = \frac{\partial J}{\partial \mathbf{x}(t)} - \mu(t) + \left[\frac{\partial \mathcal{L}(\mathbf{x}, \mathbf{\Gamma}_{prior} \mathbf{u})}{\partial \mathbf{x}(t)}\right]^T \mu(t+1) = 0 \qquad 0 < t < t_f$$

$$\frac{\partial \mathcal{L}}{\partial \mathbf{u}(t)} = \mathbf{\Gamma}_{prior}^{-1} u(t) + \left[ \frac{\partial \mathcal{L}(\mathbf{x}, \mathbf{\Gamma}_{prior} \mathbf{u})}{\partial \mathbf{u}(t)} \right] \quad \mu(t+1) = 0 \qquad 0 < t < t_f$$

Patrick Heimbach ECCO as a statistical/Bayesian inverse problem

Daa

Machine Learning		(Variational) Data Assimilation
labels	у	observations $y(t)$
features	x	state (at time $t$ ) $x(t)$
neural network	y' = W(x)	nonlinear physical $x(t+1) = \mathcal{L}(x(t))$
(surrogate model)		forward model
loss function	$J = (y - y')^2$	cost function $J = \left[ y - \mathcal{E}(\mathcal{L}(x)) \right]^T \mathbf{R}^{-1} \left[ y - \mathcal{E}(\mathcal{L}(x)) \right]$
(?)		observation ${\cal E}$
		operator
regularisation	$  \alpha  $	background $J^b = \begin{bmatrix} x_0 - x^b \end{bmatrix}^T \mathbf{B}^{-1} \begin{bmatrix} x_0 - x^b \end{bmatrix}$
stochastic gradient		Newton method
backpropagation		adjoint model $\frac{\partial J}{\partial x} = \left[ \left( \mathbf{R}^{-1} \right) \left( \frac{\partial J}{\partial y} \right) \mathbf{E} \left( \frac{\partial L}{\partial x} \right) \right]^T$
training		optimization, calibration
convolutional layers		localization
differentiable programming for AD		differentiable physics for AD

<ロ> <四> <四> <豆> <豆> <豆> <豆> <豆> = 三

DQC

Adjoint-based model calibration & state estimation

## ECCO is learning from ...

- (the most complete set of available) ocean observations
- ... AND known physics/dynamics,
- ... by solving a gigantic least-squares model-data misfit minimization
- ... using the adjoint / Lagrange Multiplier Method

#### The MIT general circulation model (MITgcm) and AD-enabled adjoint code generation

Approximated form of Navier-Stokes equations for an incompressible fluid on rotating sphere (hydrostatic or non-hydrostatic), consisting of:

- momentum equation
- conservation of mass
- conservation of heat, salt
- nonlinear equation of state for seawater
- subgrid-scale parameterizations
- scalable (domain decomposition)
- general curvilinear grid (incl. cubed-sphere)
- adjoint code generation via automatic differentiation (AD) using TAF, OpenAD



 $\int_{\sigma} \mathbf{\kappa} \times \vec{\mathbf{v}}_{h} + \frac{1}{\rho_{c}} \nabla_{z} p = \mathcal{F}$  $\epsilon_{nh} \frac{Dw}{Dt} + \frac{g\rho}{\rho_{c}} + \frac{1}{\rho_{c}} \frac{\partial p}{\partial z} = \epsilon_{nh} \mathcal{F}_{w}$ 

 $\nabla_z \cdot \vec{\mathbf{v}}_h + \frac{\partial w}{\partial z} =$ 

 $= -\rho(\theta,S)$ 

## Adjoint-based model calibration & state estimation

#### Generating & maintaining the adjoint of a state-of-the-art ocean circulation model



LECTURE NOTES IN COMPUTATIONAL SCIENCE AND ENGINEERING

Christian H. Bischof · H. Martin Bücker Paul Hovland · Uwe Naumann · Jean Utke Editors

T.Schlid

Advances in Automatic Differentiation

Springer

#### hand-written adjoint



#### Automatic Differentiation



Giering & Kaminski (1998); Marotzke et al. (1999); Heimbach et al. (2005); Utke et al. (2007); Griewank & Walther (2008)

## Filter vs. Smoother

## The virtues of propertyconserving estimation

Why adjoints: dynamical & kinematical consistency in DA

#### Numerical Weather Prediction (NWP) – a filtering problem

- Relatively abundant data sampling of the 3-dim. atmosphere
- NWP targets optimal forecasting
  - ➔ find initial conditions which produce best possible forecast;
  - $\rightarrow$  dynamical consistency or property conservation NOT required



Why adjoints: dynamical & kinematical consistency in DA

#### Numerical Weather Prediction (NWP) – a filtering problem

- Relatively abundant data sampling of the 3-dim. atmosphere
- NWP targets optimal forecasting
  - ➔ find initial conditions which produce best possible forecast;
  - $\rightarrow$  dynamical consistency or property conservation NOT required

#### Ocean state estimation/reconstruction – a smoothing problem

- Sparse data sampling of the 3-D. ocean
- Understanding past & present state of the ocean is a major goal all by itself
  - $\rightarrow$  use observations in an optimal way
  - ➔ dynamic consistency & property conservation ESSENTIAL for climate



#### Why adjoints: dynamical & kinematical consistency in DA

Balancing the momentum, freshwater, and heat budgets



Why adjoints: dynamical & kinematical consistency in DA

Tracer budgets in a global ocean reanalysis produced via filtering approach



Components in the tendency equation

dT/dt = r.h.s.

**Unphysical analysis increments** play leading role in the tracer tendencies

D. Trossman (in perpetual rejection)

Why adjoints: dynamical & kinematical consistency in DA



#### Global-ocean net heat flux imbalance

## Example Applications

Gaining insight through quantifying time-evolving property budgets

Use of observations-only vs. state estimates for understanding

Dynamics & variability of North Atlantic (Eighteen-degree) Mode Water formation



Use of observations-only vs. state estimates for understanding



The global array of Argo profiling floats http://www.argo.ucsd.edu



Use of observations-only vs. state estimates for understanding

Diabatic and adiabatic contributions to water mass volume variability in the North Atlantic subtropical gyre (26°N – 45°N) **Evans et al., JPO (2017)** 



Use of observations-only vs. state estimates for understanding

Diabatic and adiabatic contributions to water mass volume variability in the North Atlantic subtropical gyre (26°N – 45°N) **Evans et al., JPO (2017)** 



Monthly total & diathermal transformation due to air–sea heat fluxes NCEP + Reynolds SST

Monthly total & diathermal transformation due to air–sea heat fluxes ECCO v4

## Beyond optimization

Science applications with the adjoint model: sensitivity maps & dynamical attribution

#### Causal / dynamical attribution: South Atlantic Ocean Circulation

## South Atlantic Meridional Overturning Circulation (SAMOC) variability





#### Causal / dynamical attribution: South Atlantic Ocean Circulation

## South Atlantic Meridional Overturning Circulation (SAMOC) variability



Quantity of interest:

 $\delta \mathcal{J}(u(x, y, t)) \equiv Monthly AMOC Anomaly @ 34°S$ "controlled" by:

 $\delta u(x, y, t) \equiv$  Surface Atm. Forcing Perturbations

through (assumed) linear dynamics described by:

$$\frac{\partial \mathcal{J}}{\partial u}(x, y, t) \equiv \text{Sensitivity}$$

#### Smith & Heimbach, J. Clim. (2019)



#### Causal / dynamical attribution: South Atlantic Ocean Circulation

Smith & Heimbach, J. Clim. (2019)

Sensitivity of mass transport J with respect to forcings, initial conditions, carried via "backpropagation" by the time-evolving dual/adjoint state

$$\mu_{0} = \frac{\partial J}{\partial x_{0}} = \sum_{1 \le t \le t_{f}} \frac{\partial x_{t}}{\partial x_{0}} \left( \frac{\partial J}{\partial x_{t}} \right)$$
$$= \frac{\partial x_{1}}{\partial x_{0}} \left( \frac{\partial J}{\partial x_{1}} \right) + \frac{\partial x_{1}}{\partial x_{0}} \frac{\partial x_{2}}{\partial x_{1}} \left( \frac{\partial J}{\partial x_{2}} \right)$$
$$+ \dots + \frac{\partial x_{1}}{\partial x_{0}} \cdots \frac{\partial x_{t_{f}}}{\partial x_{t_{f}-1}} \left( \frac{\partial J}{\partial x_{t_{f}}} \right)$$
$$= \mathbf{L}^{T} \frac{\partial J}{\partial x_{1}} + \mathbf{L}^{T} \mathbf{L}^{T} \frac{\partial J}{\partial x_{2}} + \dots + \mathbf{L}^{T} \cdots \mathbf{L}^{T} \frac{\partial J}{\partial x_{t_{f}}}$$

 $L^T$ : is the adjoint model (and L is the tangent linear model)  $\mu_k = \left(\frac{\partial J}{\partial x_k}\right)$ : Lagrange multipliers or gradients

#### Causal / dynamical attribution: The South Atlantic Ocean Circulation

• Use of the availability of the *dual* ocean state (time-evolving adjoint state) for scientific analysis.



#### Smith & Heimbach, J. Clim. (2019)

**Causal / dynamical attribution:** The South Atlantic Ocean Circulation

Dynamic attribution of Sv interannual SAMOC variability due to wind stress perturbations via Green's functions

Smith & Heimbach, J. Clim. (2019)





# Beyond optimization

Hessian-based Uncertainty Quantification & formal observing system design What is Uncertainty Quantification in the context of Data Assimilation?

# Different uses:

- Parameter estimation (<u>calibration</u>)
   uncertainties in the parameters
- State estimation / reconstruction / synthesis (<u>interpolation</u>)
  - uncertainties in reconstructed state or derived quantities of interest (Qols)
- Forecast initialization (<u>extrapolation</u>)
   uncertainties in the forecast

Origins of uncertainty in the context of DA

#### • Observations:

- measurement error & irregular sampling (in space and time)
- Assimilation scheme:
  - DA algorithm (and approximations)
  - how observations are ingested in DA
- Model:
  - Parametric uncertainties
  - Structural uncertainties (discretization, model inadequacy)
- All boundary conditions:
  - external forcing, bathymetry, lateral boundaries
- Use of prior knowledge:
  - error covariances, representation error

Why is uncertainty quantification difficult to do?

- Spatio-temporally irregular sampling of observations
  - How to account for inhomogeneous sampling uncertainty,

## • Parametric uncertainty

- "The curse of dimensionality"
- $N^2$  for parameter space dimension N
- Structural uncertainty / model inadequacy
  - How to capture? Source terms in the model?
  - Absorb within parameter estimation?
  - Machine learning / neural networks as surrogate models?
  - Stochasticity

Away forward: Bayesian inverse methods computational frameworks that account <u>jointly</u> for model, observations, forcings, prior knowledge, and <u>all</u> their uncertainties:

- <u>MCMC and similar sampling techniques</u> hard for high-dim. parameter spaces
- <u>Derivative-based inference</u>:
  - o propagates *all* uncertainties
  - provides dynamical underpinning of uncertainty propagation
  - $\circ$  infers low-order modes

#### Approximation of the posterior PDF by posterior error covariance $\Gamma_{post}$

Posterior error covariance given by inverse of the Hessian at  $\mathbf{u}_{MAP}$ 

$$\Gamma_{post} = \left[ H_{misfit}(\mathbf{u}_{MAP}) + \Gamma_{prior}^{-1} 
ight]^{-1}$$

where

$$H_{misfit}(\mathbf{u}_{MAP}) \approx F^T \Gamma_{obs}^{-1} F$$

Gauss-Newton Hessian approximation

with:

- F: tangent linear operator of  $\mathcal{F}$
- $F^{T}$ : adjoint operator of  $\mathcal{F}$

< ロ > < 回 > < 回 > < 回 > < 回 >

#### How to obtain/extract information from $\Gamma_{post}$

• form prior-preconditioned Hessian

$$\tilde{H}_{misfit} = \Gamma_{prior} H_{misfit} ,$$

• formulate generalized eigenvalue problem

$$\tilde{H}_{misfit} W = W \Lambda$$

where

W: eigenvector matrix $r_2 = \frac{1}{\lambda_2}$  $\Lambda = \operatorname{diag}(\lambda_i)$ : eigenvalue matrix $V = \Gamma_{prior}^{1/2} W$ : prior-preconditioned eigenvector matrix



▲ □ ▶ ▲ □

< E

1

DQC

#### How to obtain/extract information from $\Gamma_{post}$

• retain only r largest eigenvalues, yields low-rank approximation:

$$\tilde{H}_{misfit} \approx V_r \Lambda_r V_r^T$$
$$\left[\tilde{H}_{misfit} + I\right]^{-1} \approx I - V_r D_r V_r^T = R_r$$

with low-rank resolution operator  $R_r$  and

$$D_r = \operatorname{diag}(\lambda_r/(\lambda_r+1))$$

Low-rank eigen-decomposition of  $\Gamma_{post}$  $\Gamma_{post} = \Gamma_{prior} - \sum_{i=1}^{r} \frac{\lambda_{i}}{\lambda_{i}+1} \left(\Gamma_{prior}^{1/2} w_{i}\right) \left(\Gamma_{prior}^{1/2} w_{i}\right)^{T}$ 

Patrick Heimbach ECCO as a statistical/Bayesian inverse problem

nar

Hessian uncertainty quantification ...

... applied to observing system design in the North Atlantic

#### **JAMES** Journal of Advances in Modeling Earth Systems

#### **RESEARCH ARTICLE**

10.1029/2020MS002386

#### **Key Points:**

- We apply Hessian uncertainty quantification (UQ) to the global ocean state estimate ECCO, and explore its use for observing system design
- Hessian UQ elucidates oceanic teleconnections that communicate observational constraints over basinscale distances
- Going beyond previous adjoint ocean modeling techniques, Hessian UQ rigorously assesses redundancy and optimality of observing systems

#### Leveraging Uncertainty Quantification to Design Ocean Climate Observing Systems

#### Nora Loose<sup>1</sup> b and Patrick Heimbach<sup>1,2,3</sup>

<sup>1</sup>Oden Institute for Computational Engineering and Sciences, The University of Texas at Austin, Austin, TX, USA, <sup>2</sup>Jackson School of Geosciences, The University of Texas at Austin, Austin, TX, USA, <sup>3</sup>Institute for Geophysics, The University of Texas at Austin, Austin, TX, USA

**Abstract** Ocean observations are expensive and difficult to collect. Designing effective ocean observing systems therefore warrants deliberate, quantitative strategies. We leverage adjoint modeling and Hessian uncertainty quantification (UQ) within the ECCO (Estimating the Circulation and Climate of the Ocean) framework to explore a new design strategy for ocean climate observing systems. Within this context, an observing system is optimal if it minimizes uncertainty in a set of investigator-defined quantities of interest (QoIs), such as oceanic transports or other key climate indices. We show that

## Predictive data science: from observation & simulation to decision



"Data assimilation is essentially an automation of the scientific method" S. Penny et al., Front. Mar. Sci. (2019)

## https://crios-ut.github.io

#### DILBERT

#### **BY SCOTT ADAMS**

